

# The TDLAS instrument for the detection of total inorganic chlorine in the stratosphere

T. J. Johnson,<sup>1</sup> G. W. Harris

Centre for Atmospheric Chemistry, York University, North York, Ontario, Canada

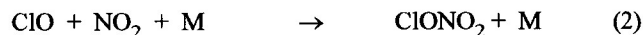
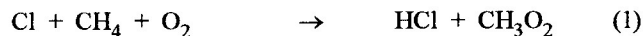
J. Bonifer, J. N. Crowley, T. Zenker,<sup>2</sup> H. Fischer

Max Planck Institute for Chemistry, Air Chemistry Department, Mainz, Germany

**Abstract.** A novel method based on tunable diode laser absorption spectroscopy is proposed for the detection of the sum of the stratospheric chlorine components HCl, ClONO<sub>2</sub>, ClOOCl and ClO. All of the above species are detected as HCl following thermal decomposition (ClONO<sub>2</sub> and ClOOCl) to ClO which is converted to Cl via reaction with NO; the Cl atom is then rapidly reacted with a hydrocarbon to form HCl. The method has been investigated in the laboratory for detection of ClONO<sub>2</sub>, for which the conversion efficiency was found to be 100% under suitable conditions.

## 1. Introduction

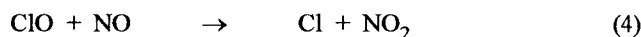
The emission of anthropogenic chlorofluorocarbons has resulted in a present-day stratospheric chlorine burden of nearly 4 ppbv (parts per billion volume mixing ratio) [World Meteorological Organization, 1992]. The observation of a significant negative trend in ozone concentrations at mid-latitude [Herman et al., 1994] and, more spectacularly, in the Antarctic [Farman et al., 1985] are closely linked to chemical reactions involving chlorinated radical species. This chemistry is initiated by the photolysis of the otherwise chemically inert chlorofluorocarbons (CFCs) to yield a chlorine atom; the Cl atom can react with O<sub>3</sub> to generate ClO (chlorine monoxide) which is the major component of the stratospheric reactive chlorine species Cl<sub>x</sub> (= Cl + Cl<sub>2</sub> + ClO + ClOOCl + OCIO + HOCl). In the normal stratosphere, most of the chlorine is found as the less reactive "reservoir species" HCl and ClONO<sub>2</sub> (Cl<sub>y</sub> = Cl<sub>x</sub> + HCl + ClONO<sub>2</sub>), which are mainly formed from Cl and ClO via the following reactions:



The stratospheric partitioning between the reactive chlorine species Cl<sub>x</sub> and the reservoir species ClONO<sub>2</sub> and HCl to a large extent determines the rate of the halogen-initiated

ozone depletion. Under "normal" conditions, inorganic chlorine is to a first approximation partitioned between the two reservoir species HCl and ClONO<sub>2</sub> in the mid-latitude stratosphere. Under "ozone hole" conditions this situation is changed and extensive activation of these chlorine reservoirs into active forms takes place. Up to 1.2 ppbv of ClO have been measured [Anderson et al., 1991], but direct *in situ* measurements of the Cl reservoir species are few. A clear requirement to better our understanding of the chemical composition of the stratosphere (and thus the ability to predict the rates of halogen-catalyzed O<sub>3</sub> depletion) is needed in the form of *in situ* measurements of total inorganic chlorine. Only very recently have direct measurements of stratospheric HCl been reported by Webster et al. [1993a, 1993b]; *in situ* concentrations of ClONO<sub>2</sub> have thus far been obtained only indirectly by a modeling study based on NO, ClO, and O<sub>3</sub> measurements [Kawa et al., 1992]. Estimates of ClONO<sub>2</sub> have also been made by comparing calculated HCl to Cl<sub>y</sub> burdens derived from measurements of speciated organic chlorine and the tracers N<sub>2</sub>O and CO<sub>2</sub> [Kawa et al., 1992, Woodbridge et al., 1995]. Specific ClONO<sub>2</sub> measurement schemes have been proposed and demonstrated as laboratory prototype experiments [Viggiano et al., 1994, Bonne et al., 1995, Perkins et al., 1995]. Other than these measurements, stratospheric ClONO<sub>2</sub> concentrations have thus far only been obtained by remote sensing experiments from either satellites [e.g. Dessler et al., 1995], the space shuttle [e.g. Zander et al., 1990], or balloon soundings [e.g. Oelhaf et al., 1994], as well as ground- and aircraft-based FTIR absorption spectroscopy [e.g. Reisinger et al., 1995, Toon et al., 1992].

The focus of the present work, therefore, is on developing a new measurement technique for stratospheric Cl<sub>y</sub> both in the form of HCl and ClONO<sub>2</sub>, as well as other Cl<sub>y</sub> components including ClO and ClOOCl. We have investigated a scheme whereby the ClONO<sub>2</sub> is thermally decomposed into ClO and NO<sub>2</sub>; the liberated ClO reacts with added NO to give Cl and NO<sub>2</sub>. Finally, the Cl radical reacts with a proton donor such as H<sub>2</sub> or an alkane to yield HCl:



The HCl produced in the final step, as well as that originally present in the air sample, is detected directly and specifically using tunable diode laser absorption spectroscopy (TDLAS),

<sup>1</sup> Now at Bruker Optics, Billerica, Massachusetts

<sup>2</sup> Now at NASA-Langley Research Center and Hampton University, Hampton, Virginia

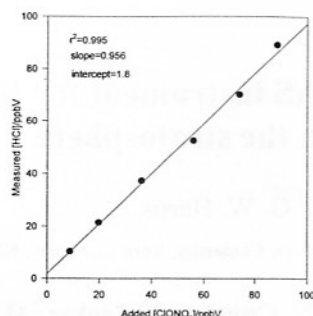
as demonstrated by Harris et al. [1992] and Webster et al. [1993b, 1994].

## 2. Experimental Methods

A ClONO<sub>2</sub> calibration source was constructed by first synthesizing a pure sample of ClONO<sub>2</sub> and using this in a dynamic dilution system. The ClONO<sub>2</sub> was synthesized by mixing Cl<sub>2</sub>O with an excess of N<sub>2</sub>O<sub>5</sub> as described previously [Davidson et al., 1987]. Substantial efforts were made to ensure purity of the final product. Broadband infrared spectra (600–4000 cm<sup>-1</sup>) of the vapor in a 10 cm cell were recorded with a Bomem DA8.02 spectrometer both at 0.5 and 0.02 cm<sup>-1</sup> resolution. Infrared spectra were monitored specifically for ClNO<sub>2</sub>, Cl<sub>2</sub>O, ClO, NO<sub>2</sub>, NO, HNO<sub>3</sub> and other impurities, until the product showed only the spectrum of pure ClONO<sub>2</sub> [Davidson et al., 1987]. To check against possible Cl<sub>2</sub> and OCIO impurities, long path (9.75 m) UV absorption measurements were undertaken; the impurity levels were found to be OCIO ~0.01% and Cl<sub>2</sub> < 1%. After all spectroscopic checks the vapor pressure of the pure compound at -78 C was measured using a 10-Torr Baratron gauge to be 1.23 Torr (1.64 mBar); we estimate the absolute purity to be ≥ 99%.

A flow of He at 1000 mBar was bubbled through the purified ClONO<sub>2</sub> held at -78 C to generate a saturated He/ClONO<sub>2</sub> flow of between 0 and 5 cm<sup>3</sup> STP min<sup>-1</sup> (sccm) (Fig. 1). In order to avoid potential decomposition of the ClONO<sub>2</sub> in a mass flow controller (MFC), the system was constructed to dilute the ClONO<sub>2</sub> flow by adding N<sub>2</sub> through MFC#2, most of which was then pumped off through MFC#3. Typically, the ~1 sccm flow from the source was diluted by 420 sccm N<sub>2</sub>, and of this only ~35 sccm was fed to the converter tube inlet, where it was further diluted in a flow of 7.5 L STP min<sup>-1</sup> (slm) zero air to yield a final mixing ratio in the approximate range 8 to 100 ppbv. All tubing prior to the converter was made of Teflon; all contact materials after the converter were of either glass or stainless steel. All mass flow controllers were multiply calibrated over the range in which they were used.

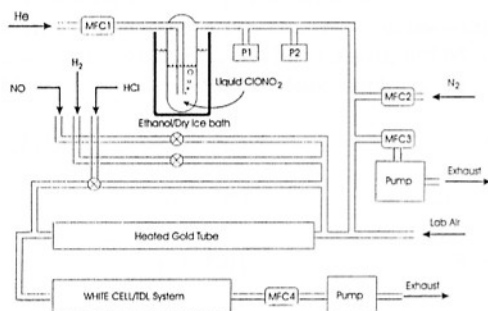
The thermal converter was a 70 cm long, 5 mm ID solid Au tube heated to 570 K over a 20 cm length near the tube center, and was originally constructed for use as a "master blaster" NO<sub>y</sub> converter. Several such NO<sub>y</sub> catalysis instruments have been described [e.g. Fahey et al., 1985]; details of the present system are described elsewhere [Zenker et al., manuscript submitted, 1996]. The converter output was coupled to the TDL White cell inlet. The TDL



**Figure 2.** Plot of TDLAS measured HCl mixing ratio versus the ClONO<sub>2</sub> mixing ratio introduced into the converter.

spectrometer used a Pb-salt laser tuned to scan the P(5) HCl line at 2775.76 cm<sup>-1</sup> and operated with 7.5 kHz modulation and standard 2f lock-in detection [Roths et al., 1996]. The White cell (Fig. 1) had a base length of 45 cm and a total optical path of 45.9 m [Wienhold et al., 1994]. The beam exiting the White cell was collected and focused by a 190 mm focal length off-axis parabolic mirror onto an HgCdTe detector. The cell was maintained at 40.6 mBar pressure with the pressure drop occurring at a calibrated glass capillary at the White cell inlet which limited the total flow to ~7.5 slm. Because of the relatively short optical path and lack of high frequency modulation techniques, the absolute detection limits for HCl were moderate in this experiment (~300 pptv); HCl detection limits an order of magnitude better can be obtained using TDLAS [Harris et al., 1992, Webster et al., 1994]. However, the present system has sufficient sensitivity to test the proposed detection method at realistic Cl<sub>y</sub> levels (approximately 5 ppbv Cl).

In order to calibrate HCl mixing ratios, either of two permeation sources were used. These operated by flowing N<sub>2</sub> gas through a Teflon tube immersed in one of two thermostatted concentrated HCl solutions. The HCl flow rates were confirmed both by optical density measurement (high concentration permeation source only) and collection of aqueous HCl samples followed by ion chromatographic (IC) analysis. A third HCl calibration source (a certified standard mixture of 500 ppbv HCl in N<sub>2</sub> - Messer Griesheim) was used to cross-calibrate the permeation sources. We performed multipoint calibrations using the high concentration permeation source (~75 ppbv HCl), the low concentration source (~8 ppbv), as well as the tank mixture in the range 2 to 15 ppbv. ClONO<sub>2</sub> flow rates were calculated using the 1.64 mBar vapor pressure of ClONO<sub>2</sub> in He along with the flow rates of the various controllers F1 - F4: reliable mixing ratios in the 10 to 120 ppbv range were obtained. The ClONO<sub>2</sub> flow rates were verified for several samples by IC analysis for nitrate ion after a timed collection of the ClONO<sub>2</sub> flow into an aqueous bubbler. For the IC measurements, several samples were collected in a 2-flask cascade fashion in order to establish collection efficiency at various flows; all IC samples had blanks subtracted. In most cases synthetic air (Messer Griesheim, alkane free, Product 0131) was used as the carrier gas entering the converter to which were added flows of NO (1.6 ppm) and H<sub>2</sub> (~1%) to obtain conversion of ClONO<sub>2</sub> to HCl. Measured HCl values were derived from least-squares fits of the TDL 2f signal to the calibration signal; reported values were corrected for drift in detector response and line centering.



**Figure 1.** Block diagram of the converter system. See text for details. MFC = Mass Flow Controller.

### 3. Results

The most important result is displayed in Figure 2's plot of measured HCl mixing ratio versus the ClONO<sub>2</sub> mixing ratio introduced into the converter. The linear and stoichiometric response of measured HCl versus input ClONO<sub>2</sub> clearly establishes the possibility of using such a system for determination of stratospheric Cl<sub>y</sub>. For each injected ClONO<sub>2</sub> flow the HCl signal was allowed to equilibrate to a reasonably stable value, approximately 10 readings (30 s averages) were recorded, and the ClONO<sub>2</sub> input withdrawn from the converter in order to collect an IC sample (requiring from 3 to 60 min), and finally the ClONO<sub>2</sub> was re-introduced to the converter. After re-equilibrating, ~10 additional values were recorded. Each point in Figure 2 thus represents the mean of the ~20 values from before and after the IC collection. The plot in Figure 2 yields a slope of 0.956 ( $\pm 0.034$ ) with an intercept of 1.8 ( $\pm 1.9$ ) and a correlation coefficient of  $r^2 = 0.995$ . If one follows the Cl atom through reactions (3) - (5), then, provided the chemistry is complete, the stoichiometry for ClONO<sub>2</sub> conversion to HCl is 1:1; hence the slope in Figure 2 is predicted to be unity. The slight positive intercept results from a memory effect which we observed whenever the inlet system and White cell had been exposed to high levels ( $\geq 50$  ppbv) of HCl. Even after ~1 hour, a residual outgassing signal equivalent to ~1.5 ppbv HCl persisted. This residual signal was not observed if only moderate levels of HCl or ClONO<sub>2</sub> (in the range 5 - 10 ppbv) were introduced into the system. In these cases the signal disappeared below the detection limit in less than 5 minutes after the calibration gas was removed.

The chemistry delineated in reactions (3) - (5) is relatively straightforward. However, for stratospheric measurements the questions of either spectroscopic interferences or unknown side reactions naturally present themselves. Brassington [1989] has shown that the HCl line we used is free of spectroscopic interference to at least the 10s of ppt level. Possible further spectroscopic interferences were investigated by replacing the bulk gas carrier flow of synthetic air (~7.5 slm) by laboratory air which is relatively rich in hydrocarbons, NO<sub>x</sub>, as well as water: No spectroscopic interferences were observed. The momentary pressure change often led to a spike in the HCl signal, but after several minutes equilibration the detected HCl returned to the same value, minus but a small fraction, typically 3%. The deviation was well within the accuracy of the HCl measurement ( $\pm 10\%$ ), and we attribute the small deviations to either slight variations in pressure and flow rates, or possibly different pressure broadening coefficients in laboratory air as compared to synthetic air. A heterogeneous reaction between ClONO<sub>2</sub> and water may also have an effect, but the measured HCl signal versus input ClONO<sub>2</sub> again showed the same linearity over the full domain of the measurement. We also note that when using either laboratory or synthetic air, the proton donor (here H<sub>2</sub>) is almost superfluous since any proton donor at the ppm level (e.g. CH<sub>4</sub> zero-air impurity) is sufficient for reaction (5) to proceed quickly. Only by using long purges with very alkane free air could we limit the final conversion step to HCl.

The converter temperature is a very important parameter, since the initial dissociation process is known to have a strong thermal dependence [Knauth, 1978]. Anderson and Fahey [1990] have investigated the rate constant for

dissociation of ClONO<sub>2</sub> in such a gold converter and obtained the value  $10^{-6.16} \exp(-90.7 \text{ kJ mol}^{-1}/RT) \text{ cm}^3 \text{ s}^{-1} \text{ molecule}^{-1}$  as measured at lower pressures (60-160 Torr). This temperature dependence is of concern in the present instance because the converter has a 20 cm copper sleeve heating element in the center with a significant temperature gradient toward either end. To investigate the temperature dependence of the conversion efficiency, a constant flow of 70.3 ppbv ClONO<sub>2</sub> was maintained with large excesses of NO and H<sub>2</sub> while monitoring the HCl signal as a function of nominal converter temperature. The experiment showed that temperatures  $\geq 550 \text{ K}$  are needed to ensure  $\geq 95\%$  dissociation and conversion to HCl.

### 4. Discussion

The present experiments validate the approach of using the reaction scheme (3) - (5) for detection of stratospheric Cl<sub>y</sub> (at least the ClONO<sub>2</sub>/HCl components) as HCl at (sub)-ppbv levels using TDL spectroscopy. The error associated with Figure 2 must incorporate uncertainties in both the x- and y-values. We estimate the total error on the ClONO<sub>2</sub> concentration in the White cell to be approximately 15%. A major source of this is the ClONO<sub>2</sub> flow dilution system, in particular the differential addition/removal scheme using MFC #2 and MFC #3, since each of these flows are only accurate to 0.25% of the full scale reading (1000 and 500 sccm, respectively), along with the smaller but non-negligible uncertainty in the ClONO<sub>2</sub> vapor pressure measurement. The total error for the HCl concentration in the White cell is estimated to be ~10%, including the error of the HCl permeation sources, the fit error, and the HCl "sticking" phenomenon during and after calibration, as discussed above. This gives a total error on the entire measurement scheme estimated at ~25%, although the regression values indicate that this estimate may be somewhat liberal. Glass and heated metals have proven to have the least memory effect for HCl. Future designs will incorporate these materials, as well as heated inlets to further reduce response times. Further known technological improvements should also help realize better detection limits. Use of an InSb detector, high frequency modulation techniques [Johnson *et al.*, 1991], and longer pathlengths will all help bring about higher sensitivity.

The present study has focused on the detection of ClONO<sub>2</sub> (and, of course, HCl) for which the method has been tested. We now discuss the detection of the other major Cl-containing stratospheric compounds ClO and ClOOCl, and the experimental modifications that would enable separation of the HCl, ClONO<sub>2</sub>, ClOOCl and ClO signals. One such variation might be the use of isotopically labeled species that could be used to differentiate the Cl<sub>y</sub> components. For example, the reaction



could be used to differentiate between HCl and [Cl<sub>y</sub>] - [HCl]. One can also envision a system with a dual inlet to distinguish between the HCl and ClONO<sub>2</sub> by using different temperatures in the inlet/converter. The thermal dissociation of ClOOCl (to 2 ClO) has recently been reinvestigated by Nickolaissen *et al.* [1994]. As the thermal decomposition rate of ClOOCl is a factor of 100 faster than for ClONO<sub>2</sub> at 600

K, we expect that  $\text{ClOOC}\text{Cl}$  will also be stoichiometrically converted to  $\text{HCl}$  in our reactor. Additionally, the thermal decomposition rate constants of  $\text{ClONO}_2$  and  $\text{ClOOC}\text{Cl}$  have different temperature dependencies and variations of temperature may provide a means to differentiate between the two. The detection of ambient  $\text{ClO}$  is ensured with this method (we have shown that  $\text{ClO}$  is 100% converted to  $\text{HCl}$ ) and, in theory, it can be selectively detected in the presence of  $\text{HCl}$ ,  $\text{ClONO}_2$ , and  $\text{ClOOC}\text{Cl}$  by simply switching the  $\text{NO}$  in and out of a cool reactor (where there is no thermal decomposition of  $\text{ClONO}_2$  or  $\text{ClOOC}\text{Cl}$ ) and measuring the signal modulation due to  $\text{HCl}$  ( $\text{NO}$  off) and  $\text{HCl} + \text{ClO}$  ( $\text{NO}$  on). Such experiments are still speculative, though, as our system has only been tested for linearity in  $\text{HCl}$  and  $\text{ClONO}_2$  inputs.

The conversion of  $\text{ClO}$  to  $\text{Cl}$  under the present conditions takes place in only 12 ms ( $\tau_{1/2} = 1/(k_{\text{ClO}+\text{NO}})[\text{NO}] = 4$  ms), while the conversion of  $\text{Cl}$  to  $\text{HCl}$  requires only 30  $\mu\text{s}$  ( $\tau_{1/2} = (1/k_{\text{Cl}+\text{H}_2})[\text{H}_2] = 10\mu\text{s}$ ). Clearly, diffusion to walls cannot compete on such short time scales, and the 100% conversion efficiency of  $\text{ClONO}_2$  to  $\text{HCl}$  is ensured. This is an advantage of the present method over those in which the  $\text{ClO}$  or  $\text{Cl}$  radicals are detected following thermal decomposition of the  $\text{ClONO}_2$  since our method avoids uncertainties as to the rate of loss of the  $\text{Cl}$  or  $\text{ClO}$  radical (via e.g. diffusion from the viewing zone or to the reactor wall).

## 5. Summary

We have demonstrated that the thermal fragmentation of  $\text{ClONO}_2$  to  $\text{ClO}$  and  $\text{NO}_2$  followed by reactions with  $\text{NO}$  and  $\text{H}_2$  quantitatively converts  $\text{ClONO}_2$  to  $\text{HCl}$ . The reaction sequence (3) - (5) has been shown to be linear and stoichiometric in  $\text{ClONO}_2$  input in the range 5 - 120 ppbv. More sensitive detection promises a system that can measure  $\text{Cl}_y$  with sufficient sensitivity and accuracy for stratospheric measurements. We are currently constructing such a system. We thank M. English for assistance in preparing the figures.

## References

- Anderson, J.G., *et al.*, Free radicals within the Antarctic vortex: The role of CFCs in Antarctic ozone loss, *Science*, **251**, 39-45, 1991.
- Anderson, L.C. and D.W. Fahey, Studies with  $\text{ClONO}_2$ : Thermal dissociation rate and catalytic conversion to  $\text{NO}$  using an  $\text{NO}/\text{O}_3$  chemiluminescence detector, *J. Phys. Chem.*, **94**, 644-652, 1990.
- Bonne, G.P. *et al.*, A laboratory prototype for *in situ* detection of  $\text{ClONO}_2$  I:  $\text{ClO}$  detection, (abstract), *Suppl. EOS*, **76**, F97, 1995.
- Brassington, D.J. in *Monitoring of gaseous pollutants by tunable diode lasers*, 1988, edited by R. Grisar, G. Schmidtke, M. Tacke, and G. Restelli, 16-24, Kluwer Academic Publ., Dordrecht, 1989.
- Davidson, J.A., *et al.*, Absolute infrared absorption cross sections for  $\text{ClONO}_2$  at 296 and 223 K, *J. Geophys. Res.*, **92**, 10,921-10,925, 1987.
- DeMore, W.B. *et al.*, Chemical Kinetics and Photochemical Data for Use in Stratospheric Modeling, Evaluation Number 11, NASA Jet Propulsion Laboratory, JPL Publ. 94-26, Pasadena CA, 1994.
- Dessler, A.E., *et al.*, Correlated observations of  $\text{HCl}$  and  $\text{ClONO}_2$  from UARS and implications for stratospheric chlorine partitioning, *Geophys. Res. Lett.*, **22**, 1721-1724, 1995.
- Fahey, D.W., *et al.*, Evaluation of a catalytic reduction technique for the measurement of total reactive odd-nitrogen  $\text{NO}_y$  in the atmosphere, *J. Atmos. Chem.*, **3**, 435-468, 1985.
- Farman, J. C., *et al.*, Large losses of total ozone in Antarctica reveal seasonal  $\text{ClO}_x/\text{NO}_x$  interaction, *Nature*, **315**, 207-210, 1985.
- Harris G.W., D. Klemp and T. Zenker, An upper limit on the  $\text{HCl}$  near-surface mixing ratio at  $28^\circ\text{N}$ ,  $30^\circ\text{W}$  measured using TDLAS, *J. Atmos. Chem.*, **15**, 327-332, 1992.
- Herman, J.R. and D. Larko, Low ozone amounts during 1992-1993 from Nimbus 7 and Meteor 3 total ozone mapping spectrometers, *J. Geophys. Res.*, **99**, 3483-3496, 1994.
- Johnson, T.J., *et al.*, Frequency modulation spectroscopy at 1.3  $\mu\text{m}$  using InGaAsP lasers: a prototype field instrument for atmospheric chemistry research, *Appl. Opt.*, **30**, 407-413, 1991.
- Kawa, S.R., *et al.*, Photochemical partitioning of the reactive nitrogen and chlorine reservoirs in the high-latitude stratosphere, *J. Geophys. Res.*, **97**, 7905-7923, 1992.
- Knauth, H.-D., Über den thermischen Zerfall von  $\text{ClONO}_2$  in Gegenwart von  $\text{NO}$ ,  $\text{ClNO}$  und  $\text{N}_2$ , *Ber. Bunseng. Phys. Chem.*, **82**, 212-216, 1978.
- Nickolaisen, S.L., *et al.*, Kinetics and mechanism of the  $\text{ClO} + \text{ClO}$  reaction: Pressure and temperature dependencies of the bimolecular and termolecular channels & thermal decomposition of chlorine peroxide, *J. Phys. Chem.*, **98**, 155-169, 1994.
- Oelhaf, H., *et al.*, Stratospheric  $\text{ClONO}_2$  and  $\text{HNO}_3$  profiles inside the Arctic vortex from MIPAS-B limb emission spectra obtained during EASOE, *Geophys. Res. Lett.*, **21**, 1263-1266, 1994.
- Perkins, K.K., *et al.*, A laboratory prototype for *in situ* detection of  $\text{ClONO}_2$  II:  $\text{NO}_2$  Detection, (abst.), *Suppl. EOS*, **76**, F97, 1995.
- Reisinger, A.R. *et al.*, Southern hemisphere midlatitude ground-based measurements of  $\text{ClONO}_2$ : Method of analysis, seasonal cycle and long-term trend, *J. Geophys. Res.*, **100**, 23183-23193, 1995.
- Roths, J., *et al.*, FLAIR, A four-laser airborne infrared spectrometer for atmospheric trace gas measurements, accepted for publication, *Appl. Opt.*, 1996.
- Toon, O.B., *et al.*, Composition measurements of the 1989 Arctic winter stratosphere by airborne infrared solar absorption spectroscopy, *J. Geophys. Res.*, **97**, 7939-7961, 1992.
- Viggiano, A.A., R.A. Morris and J.M. Van Doren, Ion Chemistry of  $\text{ClONO}_2$  involving  $\text{NO}_3^-$  core ions: A detection scheme for  $\text{ClONO}_2$  in the atmosphere, *J. Geophys. Res. D*, **99**, 8221-8224, 1994.
- Webster, C.R. *et al.*, Chlorine chemistry on polar stratospheric cloud particles in the Arctic winter, *Science*, **261**, 1130-1134, 1993a.
- Webster, C.R., *et al.*, *In situ* measurements of the  $\text{ClO}/\text{HCl}$  ratio: Heterogeneous processing on sulfate aerosol and polar stratospheric clouds, *Geophys. Res. Lett.*, **20**, 2523-2526, 1993b.
- Webster, C.R., *et al.*, Aircraft (ER-2) laser infrared absorption spectrometer (ALIAS) for *in situ* stratospheric measurements of  $\text{HCl}$ ,  $\text{N}_2\text{O}$ ,  $\text{CH}_4$ ,  $\text{NO}_2$ , and  $\text{HNO}_3$ , *Appl. Opt.*, **33**, 454, 1994.
- Wienhold, F.G. *et al.*, Measurement of  $\text{N}_2\text{O}$  fluxes from fertilized grassland using a fast response tunable diode laser spectrometer, *J. Geophys. Res.*, **99**, 16,557-16,567, 1994.
- Woodbridge, E.L., *et al.*, Estimates of total organic chlorine in the lower stratosphere from *in situ* and flask measurements during AASE II, *J. Geophys. Res.*, **100**, 3057-3064, 1995.
- World Meteorology Org., Scientific assessment of stratospheric ozone: Rep. 25, *Global ozone res. and monit.*, Geneva, 1992.
- Zander, R., *et al.*, Stratospheric  $\text{ClONO}_2$ ,  $\text{HCl}$ , and  $\text{HF}$  concentration profiles derived from atmospheric trace molecule spectroscopy experiment spacelab 3 observations: an update *J. Geophys. Res.*, **20**, 519-20,525 1990.

T.J. Johnson and G.W. Harris, Centre for Atmospheric Chemistry, York University, 4700 Keele Street, North York, Ontario M3J 1P3 Canada

J. Bonifer, J.N. Crowley, T. Zenker and H. Fischer, Max Planck Institute for Chemistry, Air Chemistry Department, P.O. Box 3060, D-55020, Mainz, FRG.

(Received: July 30, 1996; revised October 10, 1996; accepted October 18, 1996)