

Flow reactor studies of the stable carbon isotope composition of secondary particulate organic matter generated by OH-radical-induced reactions of toluene

Satoshi Irei^a, Lin Huang^b, Fabrice Collin^{a,1}, Wendy Zhang^b,
Donald Hastie^a, Jochen Rudolph^{a,*}

Abstract

Secondary particulate organic matter (POM) formed in a flow reactor by the OH-radical-induced reactions of toluene was collected on quartz fiber filters, and its stable carbon isotope ratio was analyzed by off-line combustion and subsequent dual-inlet isotope ratio mass spectrometry. The toluene consumption in these experiments ranged from 7% to 29%. The stable carbon isotope composition ($\delta^{13}\text{C}$) of the secondary POM was in the range of -32.2‰ to -32.9‰ (VPDB scale), with some indication for a slight dependence on the extent of toluene consumption. These measured $\delta^{13}\text{C}$ values were, on average, 5.8‰ lighter than those of the parent toluene. Those observations are slightly lower than the $\delta^{13}\text{C}$ values of the sum of all toluene oxidation products (from -31.6‰ to -32.3‰) that are predicted using the kinetic isotope effect for the reaction of toluene with OH-radical under these reaction conditions and the initial $\delta^{13}\text{C}$ of the parent toluene. Therefore, mass balance dictates that the fractionation between gas-phase and particle-phase products is small. On average, the particle-phase products are $0.6\text{‰} \pm 0.2\text{‰}$ lighter than the gas-phase products. This is in agreement with the concept that the initial reaction of toluene with the OH-radical is the slowest step in the reaction sequence resulting in POM formation.

Keywords: Stable carbon isotope; Kinetic isotope effect; Toluene oxidation; Secondary organic aerosols; Aerosol yield

1. Introduction

There is strong evidence that atmospheric particulate matter plays a major role in climate change (Anderson et al., 2003; Ramanathan et al., 2001) and can also have substantial negative impact on human health (Thurston et al., 1994; Urch et al., 2004). Although particulate organic matter (POM) comprises 17–44% of the mass of atmospheric particulate matter (Blanchard et al., 2002), its

chemical composition has only been partly characterized. In particular, the contribution of secondary POM is poorly understood. One of the newly emerging methods that are used to study the origin and atmospheric chemistry of airborne POM is the measurement of stable carbon isotope composition ($\delta^{13}\text{C}$). Effectively, all published studies on the isotopic composition of POM target the origin of primary POM (Chesselet et al., 1981; Kaplan and Gordon, 1994; Simoneit, 1997; Norman et al., 1999; Narukawa et al., 1999; Okuda et al., 2002; Conte and Weber, 2002). To the best of our knowledge, no study on the $\delta^{13}\text{C}$ for secondary POM has been reported to date. Based on the observation of substantial isotope fractionation associated with the most important atmospheric reactions of non-methane hydrocarbons (Rudolph et al., 2000), it seems likely that the photochemical products, including secondary POM, differ in $\delta^{13}\text{C}$ from the precursors. The studies on $\delta^{13}\text{C}$ for formic and acetic acids in the collected rainwater in Los Angeles (Sakugawa and Kaplan, 1995) support this expectation. One piece of information missing in a stable carbon isotope study for formation processes of secondary POM in the atmosphere is the isotope fractionation during the chemical reactions that result in POM formation.

In this paper, we present laboratory studies on the $\delta^{13}\text{C}$ values of secondary POM formed by the reaction of toluene with the OH-radical. The reaction with the OH-radical was chosen because it is by far the most important atmospheric removal

process for hydrocarbons. Toluene was selected since secondary POM formation from this precursor is well established (Izumi and Fukuyama, 1990; Wang et al., 1992; Forstner et al., 1997; Hurley et al., 2001; Kleindienst et al., 2004; Sato et al., 2004; Stroud et al., 2004). Furthermore, the kinetic isotope effect (KIE) for the reaction of toluene with the OH-radical is well known (Anderson et al., 2004). This allows the comparison of the isotope fractionation due to the initial step of the toluene oxidation reaction with the overall fractionation for formation of secondary POM.

2. Experiment

Secondary POM was generated in a continuous flow reactor (Fig. 1) by irradiating air containing approximately $40\ \mu\text{mol mol}^{-1}$ toluene, $230\ \mu\text{mol mol}^{-1}$ isopropyl nitrite (IPN) and $5\ \mu\text{mol mol}^{-1}$ nitrogen oxide (BOC Gases) with a 1000 W Xenon Arc lamp (Model C-50, Oriol Optical Corp.). IPN was synthesized following the procedure described by Noyes (1943) for the synthesis of *n*-butyl nitrite. OH-radicals were produced via photolysis of IPN in the presence of oxygen and nitric oxide. The efficiency of the toluene oxidation in the reactor was determined by measuring the toluene concentration at the end of the reactor with and without irradiation. These measurements were made by collecting a gaseous sample with a $250\ \mu\text{l}$ gastight syringe and analyzing the toluene content of these samples by GC-FID (HP5890, Agilent Technology).

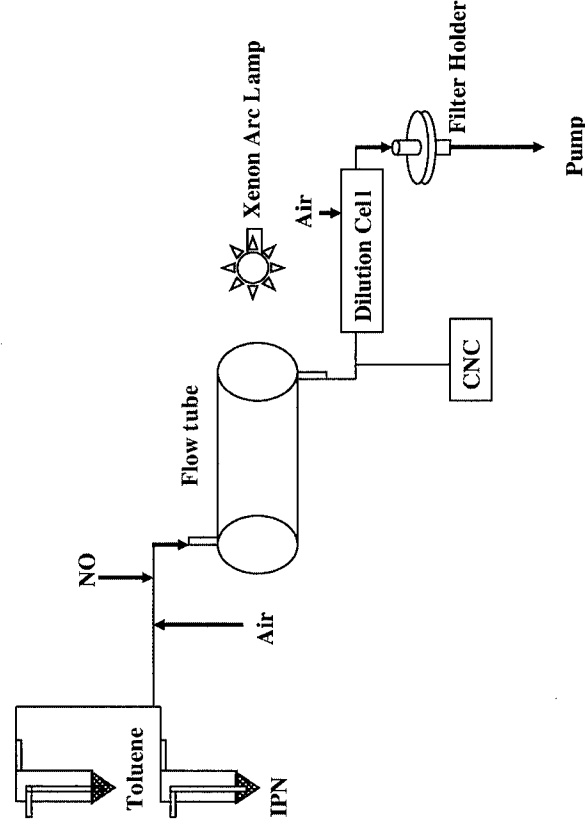


Fig. 1. Schematic diagram for the flow reactor experiment; see text for detailed explanations.

For the first set of experiments (Series A), there was no absolute calibration for toluene. Therefore, only the relative change in concentration could be derived from these measurements. For the other experiments, in which the position of the lens focussing the light source was varied in order to change the photon flux density in the flow reactor, the toluene concentrations were determined by comparison with a commercial $\mu\text{mol mol}^{-1}$ level hydrocarbon standard with known concentration. The precision of these measurements and the accuracy of the standard are better than 3% and 10%, respectively.

The total flow rate through the reactor was $0.78 \text{ dm}^3 \text{ min}^{-1}$. Two reactors of different sizes with volumes of 2.5 dm^3 (for series A, C and D) or 0.47 dm^3 (for series B) were used. POM formed in the reactor was collected on quartz fiber filters (Pallflex, Tissuequartz 2500QAT-UP 25 mm disks, PallGelman Sciences) at the outlet of the flow reactor. Forty percent of the flow through the reactor was used to monitor the formation of the secondary POM with a condensation nucleus counter (CNC; Model 3020, TSI) at the outlet of the reactor. In experiments of series A, the particle concentrations were monitored not only prior to the sampling of POM but also during POM collection. Therefore in these experiments, only 60% of the total POM formed was collected on the filters. In experiments from series B–D, the connection line to the CNC was closed after the CNC showed stable particle number concentrations, and all the flow from the reactor outlet passed through the quartz fiber filters. All quartz filters were cleaned prior to use by baking at 1123 K in an atmosphere of synthetic air. Typically double filters were used to collect the POM, and we will refer to the first filter as the particulate matter sample and to the second filter as the back-up filter. The duration of sample collection varied between 30 min and 46 h.

The samples were placed together with approximately 1 g of copper oxide, which served as the source of oxygen, in vacuum-sealed quartz tubes and combusted at 1173 K for 12 h. The amount of CO_2 in each sample was determined using an absolute pressure transducer (Intelligent Transmitter Series 6000, Paroscientific Inc.) attached to a glass bulb of known volume. This transducer simultaneously measures temperature and the CO_2 pressure, which were used to determine the CO_2 yield. The CO_2 was then transferred to a Pyrex ampoule. The $\delta^{13}\text{C}$ values of these samples were

measured with an isotope ratio mass spectrometer, (Isoprime, GV-Instruments) with a dual-inlet system. Samples with sufficient CO_2 mass were analyzed using the conventional dual bellows method, whereas small samples of a few micromoles or less were analyzed using a cold finger (i.e., a cryogenic trap) technique. For both methods the precision of the measurements was 0.02‰. However, for measurements made at different intensities for sample and standard signals, corrections had to be made due to the limited linearity of the instrument. For mass 44 signal intensities exceeding 5 nA, the linearity corrections are less than 0.7‰ and, therefore, the uncertainty resulting from these corrections is small. We estimate that under these conditions the uncertainty of the $\delta^{13}\text{C}$ measurements is less than 0.1‰. For the few measurements made with mass 44 signal intensities of less than 5 nA, the deviation from the linearity resulted in uncertainties that may be as high as 0.2–0.3‰.

All $\delta^{13}\text{C}$ measurements were made relative to a standard traceable to Vienna Pee Dee Belemnite (VPDB) scale; details of the traceability have been discussed elsewhere (Huang et al., 2003). The $\delta^{13}\text{C}$ values presented here are expressed as δ value in permille (i.e., $\delta^{13}\text{C}_{\text{sample}} = [R_{\text{sample}}/R_{\text{std}} - 1] \times 1000\text{‰}$) relative to VPDB, where R_{sample} and R_{std} represent the atomic ratios of $^{13}\text{C}/^{12}\text{C}$ in the sample and the standard, respectively. The stated uncertainties for the measured $\delta^{13}\text{C}$ and the carbon yields are the standard errors, unless otherwise noted.

Blanks and back-up filters were analyzed using a procedure identical to that used for POM samples. A background test was carried out by pumping the gas mixture through a sampling filter for 30.7 h without UV irradiation in order to evaluate adsorption of impurities and gas-phase organic material on the filters. For series A experiments, several back-up filters were combined into one sample since the amount of carbon on these filters was too low for $\delta^{13}\text{C}$ measurements by the conventional dual-inlet technique.

In order to evaluate the accuracy of the experimental procedure, filters spiked with a laboratory sucrose standard, IAEA-CH-6 (sucrose), or USGS 24 (graphite) were analyzed using the procedure described above. Two different combustion temperatures, 853 and 1173 K, were used in these tests.

The $\delta^{13}\text{C}$ of the toluene used in the flow reactor experiments was determined using a method very similar to that used for the POM samples. These samples were combusted at 853 K. Two types of

toluene samples, liquid toluene and evaporated toluene, were analyzed in order to evaluate the isotopic composition of the precursor. The liquid toluene samples were taken directly from the bulk liquid phase. The evaporated toluene samples were collected by condensation at 201 K from a gas flow saturated with toluene at 273 K.

3. Results and discussion

3.1. Measurements of blank, reference material and back-up filter

The average blank value of cleaned filters from 10 measurements is 1.0 $\mu\text{mol C}$, with a standard deviation of 0.5 $\mu\text{mol C}$ and a standard error of 0.2 $\mu\text{mol C}$. The average $\delta^{13}\text{C}$ value of the filter blanks was -22.8% with a standard deviation of 2.0‰ and a standard error of 0.6‰.

The results of reference material measurements are summarized in Table 1 together with the recommended $\delta^{13}\text{C}$ value. Note that the $\delta^{13}\text{C}$ values shown in the table are obtained from Keeling plots (see below). The results for the laboratory sucrose reference show that the recoveries for the samples combusted at 853 K are slightly lower than those combusted at 1173 K. Furthermore, evidence for incomplete combustion such as gray spots on the filters could be seen for samples oxidized at 853 K, whereas no such indicators were found for samples

oxidized at 1173 K. Nevertheless, the measured $\delta^{13}\text{C}$ values themselves have a standard deviation of less than 0.1‰ and show very good agreement with the reference value of $\delta^{13}\text{C}$ for our laboratory sucrose standard, independent of the oxidation temperature.

The tests performed with IAEA-CH-6 covered a wide range of carbon mass from approximately 1 μmol to over 30 μmol , with the lower end very close to that found on blank filters. The $\delta^{13}\text{C}$ values measured for these samples range from -11% to approximately -17% , with a clear dependence on carbon mass. This strongly suggests that this variability in the $\delta^{13}\text{C}$ values is due to the influence of the filter blanks. Indeed, the Keeling plot, which is a plot of measured $\delta^{13}\text{C}$ values versus the reciprocal of the carbon mass, gives a straight line with a very good regression coefficient (Fig. 2). Such plots can be used to obtain the blank-corrected $\delta^{13}\text{C}$ of IAEA-CH-6; the y -intercepts of the linear regressions reflect the $\delta^{13}\text{C}$ of samples with an indefinitely large amount of carbon, and therefore no influence of the blank. The intercepts obtained are $-10.4\% \pm 0.4\%$ ($R^2 = 0.982$) and $-10.3\% \pm 0.2\%$ ($R^2 = 0.994$), for combustion at 853 and 1173 K, respectively. These values completely agree with the IAEA-recommended value of $-10.4\% \pm 0.2\%$. Similarly, a high correlation with an intercept of $-15.9\% \pm 0.1\%$ ($R^2 = 0.927$) was observed for USGS 24, which was combusted at 1173 K: the $\delta^{13}\text{C}$ value recommended by IAEA is $-16.0\% \pm 0.1\%$. However, combustion of graphite at 853 K was not only incomplete, but also resulted in large uncertainties in the $\delta^{13}\text{C}$ values (Table 1). Therefore, at this combustion temperature, our

Table 1
Results of measurements of reference materials

	<i>n</i>	Spiked amount ($\mu\text{mol C}$) ^a	Recovery (%) ^{a,b}	$\delta^{13}\text{C}_{\text{VPDB}}$ (‰) ^c
Sucrose (853 K)	3	19–47	89–98	-11.65 ± 0.01
IAEA-CH-6 (853 K)	4	0.92–37	66–96	-10.4 ± 0.4
USGS 24 (853 K)	3	14–28	14–30	-17.8 ± 11
Sucrose (1173 K)	3	19–47	92–99	-12.0 ± 0.2
IAEA-CH-6 (1173 K)	4	0.92–37	88–94	-10.3 ± 0.2
USGS 24 (1173 K)	3	14–24	72–91	-15.9 ± 0.1
<i>Reference/recommended values</i>				
Sucrose ^d	—	—	—	-12.0 ± 0.1
IAEA-CH-6 ^e	—	—	—	-10.4 ± 0.2
USGS 24 ^e	—	—	—	-16.0 ± 0.1

^aRange of minimum–maximum value.

^bBlank correction was made.

^cAverage \pm standard error determined from Keeling plots.

^dOur laboratory reference value \pm standard deviation.

^eRecommended values given by IAEA.

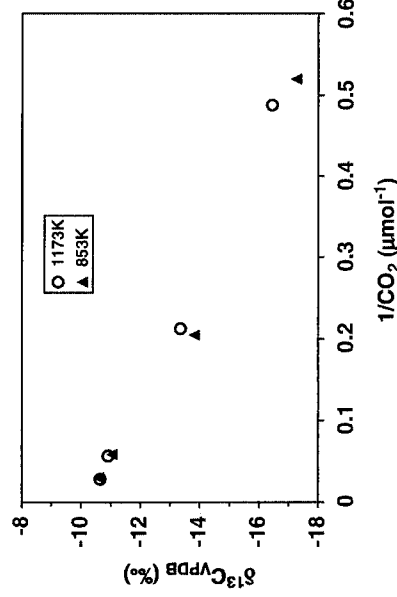


Fig. 2. Plot of isotope ratio as a function of reciprocal of CO_2 amount (Keeling plot) for IAEA-CH-6.

method is not suitable for materials that are difficult to oxidize, such as graphite.

The results of the measurements of individual back-up filters are summarized in Table 2. The combustion of the filter from the background test gave a total of 2.6 $\mu\text{mol C}$, which is 1.6 $\mu\text{mol CO}_2$ higher than that from the filter blanks. The combustion of seven back-up filters from the series A experiments combined into one sample gives a total of 12.0 $\mu\text{mol CO}_2$ and a $\delta^{13}\text{C}$ value of -28.2% . The measurement results for individual back-up filters from series B–D show that in spite of the

considerable variability in the carbon mass and the $\delta^{13}\text{C}$ found in the analysis of the back-up filters, on average the mass on the back-up filters is $1.7 \pm 0.3 \mu\text{mol C}$ higher than the blank values. On average, this difference is statistically significant. Similarly, there is a highly significant difference in $\delta^{13}\text{C}$ ($4.3\% \pm 0.7\%$ lighter than $\delta^{13}\text{C}$ of blank value). Based on mass balance considerations, the $\delta^{13}\text{C}$ value of the additional 1.7 $\mu\text{mol C}$ on the back-up filters has to be on average $-29.6\% \pm 1.0\%$, which is $6.8\% \pm 1.0\%$ lighter than the carbon on blank filters. This value is somewhat lighter than the $\delta^{13}\text{C}$ value of the toluene in the reactor and thus indicates that the additional carbon on the back-up filters is at least partly from toluene oxidation products that have $\delta^{13}\text{C}$ values of approximately -32% (see below).

There are different possibilities for the origin of carbon in excess of the blank values on the back-up filters: (i) the partial breakthrough of very small particles through the front filter, (ii) the irreversible adsorption of gas-phase oxidation products of toluene and (iii) a reversible adsorption of organic compounds in equilibrium with the gas phase. For the first two possibilities, we expect that the mass of carbon on the back-up filter increases with sampling time. Consequently, the $\delta^{13}\text{C}$ of total carbon on the back-up filters will decrease with increasing sampling duration, since the $\delta^{13}\text{C}$ of the toluene oxidation products is lower than that of the filter blanks. For the last possibility, we expect no dependence between carbon mass or $\delta^{13}\text{C}$ and sampling time. The results from linear regressions of $\delta^{13}\text{C}$ and carbon mass on the back-up filters versus experiment duration show some evidence of systematic time dependencies (Table 3). However, the slopes are very small and their uncertainties are substantial. Moreover, the time-independent parameters derived from the linear regressions are not

Table 2
Results of back-up filter measurements

Experiment	Sampling duration (h)	CO ₂ yield (μmol)	$\delta^{13}\text{C}_{\text{VPDB}}$ (%)
<i>Series B</i>			
B-1 back-up	7.4	0.8	-25.4
B-2 back-up	14.1	2.5	-26.5
B-3 back-up	24.1	2.6	-26.7
B-4 back-up	45.3	2.9	-27.5
Average B	2.2	2.2	-26.5
Standard deviation	0.9	0.9	0.9
Standard error	0.5	0.5	0.5
<i>Series C</i>			
C-1 back-up	5.1	3.3	-27.4
C-2 back-up	3.0	2.8	-27.3
C-3 back-up	15.6	3.6	-28.5
C-4 back-up	8.1	2.9	-27.6
C-5 back-up	16.8	4.1	-28.4
C-6 back-up	3.4	2.5	-26.8
Average C	3.2	3.2	-27.7
Standard deviation	0.6	0.6	0.7
Standard error	0.3	0.3	0.3
<i>Series D</i>			
D-1 back-up	2.8	1.1	-27.1
D-2 back-up	1.5	2.6	-25.7
D-3 back-up	3.2	2.9	-27.0
D-4 back-up	5.4	3.5	-27.5
D-5 back-up	0.73	2.8	-26.4
Average D	2.7	2.7	-26.8
Standard deviation	0.9	0.9	0.7
Standard error	0.4	0.4	0.3
Average all	2.7	2.7	-27.1
Standard deviation	0.9	0.9	0.9
Standard error	0.2	0.2	0.2

Table 3

Dependence between duration of sampling and mass of carbon or stable carbon isotope ratio on back-up filters

Experiment	Mass carbon		Isotope ratio	
	Slope ($\mu\text{mol C h}^{-1}$)	Intercept ($\mu\text{mol C}$)	Slope ($\% \text{h}^{-1}$)	Intercept (%)
Series B	0.04 ± 0.03	1.2 ± 0.7	-0.05 ± 0.01	-25.4 ± 0.4
Series C	0.09 ± 0.02	2.5 ± 0.2	-0.10 ± 0.02	-26.7 ± 0.2
Series D	0.14 ± 0.28	2.2 ± 0.9	-0.32 ± 0.13	-25.9 ± 0.4

identical to the values obtained from analysis of filter blanks. These differences in $\delta^{13}\text{C}$ are statistically highly significant. Thus it seems very likely that the difference in carbon mass and $\delta^{13}\text{C}$ between blank and back-up filters is due to time-dependent (i.e., sampling duration) as well as time-independent processes.

3.2. Isotopic compositions of toluene and secondary POM and POM yields

The average $\delta^{13}\text{C}$ of toluene taken directly from the bulk supply was $-27.05\% \pm 0.02\%$ ($n = 3$). The average $\delta^{13}\text{C}$ of toluene condensed from air saturated with toluene at 273 K over 4 h was $-26.72\% \pm 0.01\%$ ($n = 6$). This small, but significant difference of $0.33\% \pm 0.02\%$ agrees with the observation by Harrington et al. (1999) of a small inverse isotope fractionation of toluene during evaporation.

The carbon amount in the POM collected on the filters ranges from 2 to $94 \mu\text{mol C}$ (Table 4). Furthermore, within each series of experiments, it was proportional to the sampling time. The extent of photochemical processing of toluene (i.e., the relative change in concentrations measured without and with UV irradiation) ranges from 7% to 29% for the different experiment series. The average rates of the secondary POM collection were between 0.16 and $6.4 \mu\text{mol Ch}^{-1}$. Within each experiment series, the POM collection rates showed little variability, indicating that the reaction condition in the flow reactor remained constant throughout each series of experiments. This is also supported by Keeling plots from the different experiment series, which demonstrate that most of the variability in the $\delta^{13}\text{C}$ within a given series of experiments can be explained by the impact of the blank on the total mass of carbon. That is, the carbon mass and $\delta^{13}\text{C}$ of the blank were consistent within each experiment.

The average blank-corrected $\delta^{13}\text{C}$ values derived from Keeling plots and their standard errors are included in Table 4. They range from approximately -32.2% to -32.9% . It should be noted that for series C experiments, the data point at -31.93% significantly deviates from the linear regression line. Removing this data point changes the result of the Keeling plot from -32.57% to -32.69% . This change of 0.1% is small compared to the standard error of 0.2%. It should be noted that using Keeling plots to derive the blank-corrected $\delta^{13}\text{C}$ values also removes the impact of all time-independent bias

Table 4
Results of secondary POM analysis

Experiment	Sampling duration (h)	CO ₂ yield ^a (μmol)	POM carbon formation rate ($\mu\text{mol Ch}^{-1}$) ^b	$\delta^{13}\text{C}_{\text{VPDB}}$ ($\%$)
<i>Series A</i>				
A-1	9.3	58.3	6.2	-32.07
A-2	3.1	20.1	6.2	-31.38
A-3	6.3	44.0	6.8	-31.89
A-4	9	53.0	5.8	-32.06
A-5	5.5	40.7	7.2	-31.91
A-6	3.4	22.8	6.4	-31.68
All			6.4 ^c	-32.37 ^d
Standard error			0.2	0.08
<i>Series B</i>				
B-1	7.4	2.5	0.20	-30.63
B-2	14.1	3.4	0.17	-31.14
B-3	24.1	5.1	0.17	-31.81
B-4	45.3	6.2	0.12	-31.95
All			0.16 ^c	-32.90 ^d
Standard error			0.02	0.09
<i>Series C</i>				
C-1	5.1	40.7	7.8	-31.93
C-2	3.0	16.1	5.1	-31.99
C-3	15.6	94.1	6.0	-32.54
C-4	8.1	37.0	4.4	-32.36
C-5	16.8	70.6	4.1	-32.58
C-6	3.4	15.8	4.4	—
All			5.3 ^c	-32.57 ^d
Standard error			0.6	0.20
<i>Series D</i>				
D-1	2.8	20.1	6.8	-31.94
D-2	1.5	9.8	5.9	-31.74
D-3	3.2	20.1	6.0	-31.99
D-4	5.4	32.0	5.7	-32.14
D-5	0.73	4.4	4.6	-31.30
All			5.8 ^c	-32.19 ^d
Standard error			0.4	0.04

Relative decrease in toluene concentrations for series A-D are $15\% \pm 3\%$, $7\% \pm 1\%$, $20\% \pm 2\%$ and $29\% \pm 2\%$, respectively.

^aNo blank correction made.

^bCorrected for filter blank values.

^cAverage of POM formation rate.

^d $\delta^{13}\text{C}$ values and their standard errors determined from Keeling plots.

such as adsorption of VOC on the filter in equilibrium with the gas phase. However, the procedure does not correct for potential bias from irreversible adsorption of gas-phase VOC. From the back-up filter analysis (see above), the impact of time-dependent bias is in the range of

0.1 $\mu\text{mol Ch}^{-1}$. Except for series B experiments, this is a very small effect compared to the collection rate of POM. As the analysis of the back-up filters furthermore indicates that the $\delta^{13}\text{C}$ of the additional carbon is only slightly different to that of the secondary POM, we expect that the potential bias from adsorption of gas-phase VOC on the filter to the $\delta^{13}\text{C}$ of POM derived from experiment series A, C, and D is insignificant.

For experiment series B, the situation is not so simple. Here the POM collection rate is only $0.16 \pm 0.02 \mu\text{mol Ch}^{-1}$. From the back-up filter analysis for series B experiments, the potential bias is $0.04 \pm 0.03 \mu\text{mol Ch}^{-1}$, which is a significant fraction of the total carbon collection rate for series B. The $\delta^{13}\text{C}$ of the additional carbon on the back-up filters is estimated to be $-29.6\% \pm 1.0\%$. The resulting potential bias for series B experiments therefore may be as high as 0.8% , however with an uncertainty of $\pm 0.7\%$. In view of the large uncertainty of the potential bias, it does not seem very useful to correct the results of series B experiments. However, it has to be remembered that the $\delta^{13}\text{C}$ of POM derived from Keeling plots may be somewhat heavier than the actual $\delta^{13}\text{C}$ because of the additional carbon from the gas-phase toluene oxidation products or other time-dependent artifacts.

The toluene mixing ratio in the reactor without irradiation was measured to be on average $38 \pm 7 \mu\text{mol mol}^{-1}$. The total flow rate through the reactor was 0.78 L min^{-1} , thus the total mass flow of toluene is equivalent to $509 \pm 93 \mu\text{mol Ch}^{-1}$. The percentage of photo-oxidation of toluene from the series A–D experiments was $15\% \pm 3\%$, $7\% \pm 1\%$, $20\% \pm 2\%$ and $29\% \pm 2\%$, which corresponds to the oxidation of 76 ± 14 , 36 ± 7 , 102 ± 19 and $148 \pm 27 \mu\text{mol Ch}^{-1}$, respectively. A comparison with the observed formation rates of particulate carbon (Table 5) gives particulate carbon yields of $8\% \pm 2\%$, $0.4\% \pm 0.1\%$, $5\% \pm 1\%$ and $4\% \pm 1\%$, respectively.

3.3. Dependence between the isotopic composition of secondary POM and the toluene KIE

The difference between the gas-phase toluene and particulate matter formed by the oxidation of toluene ranges from 5.5% to 6.2% , with some indications for a systematic dependence on the extent of photochemical processing of toluene. The flow reactor can be treated as a system with a closed mass balance and thus the dependence between $\delta^{13}\text{C}$ and concentration of toluene at time t can be expressed by a Rayleigh-type function:

$$\delta^{13}\text{C}_{\text{Tot}} = \left(\delta^{13}\text{C}_{\text{Tot}} + 1 \right) \left(\frac{c_{\text{Tot}}}{\theta c_{\text{Tot}}} \right)^{\frac{(-\epsilon_{\text{Tot}}^{\text{OH}}/\epsilon_{\text{Tot}})/(1+\epsilon_{\text{Tot}}^{\text{OH}})}{\theta}} - 1. \quad (1)$$

Here, $\delta^{13}\text{C}_{\text{Tot}}$ and $\delta^{13}\text{C}_{\text{Tot}}$ are the $\delta^{13}\text{C}$ of toluene at time t and at $t = 0$, respectively, and c_{Tot} and θc_{Tot} the corresponding concentrations. $\epsilon_{\text{Tot}}^{\text{OH}}$ is the KIE for the reaction of toluene with OH-radicals (i.e., $\epsilon_{\text{Tot}}^{\text{OH}} = k_{12}k_{13}^{-1}$), which is $5.95\% \pm 0.28\%$ (Anderson et al., 2004). It should be noted that Eq. (1) does not explicitly contain the time t , but c_{Tot} depends, among other factors, on processing time and thus Eq. (1) implicitly includes the time dependence of $\delta^{13}\text{C}_{\text{Tot}}$. The $\delta^{13}\text{C}$ of the sum of all oxidation products ($\delta^{13}\text{C}_{\text{Allprod}}$) accumulated at time t can be expressed by the following equation:

$$\delta^{13}\text{C}_{\text{Allprod}} = \frac{\delta^{13}\text{C}_{\text{Tot}} + 1}{1 - (c_{\text{Tot}}/\theta c_{\text{Tot}})} \left[1 - \left(\frac{c_{\text{Tot}}}{\theta c_{\text{Tot}}} \right)^{\frac{1/(1+\epsilon_{\text{Tot}}^{\text{OH}})}{\theta}} \right] - 1. \quad (2)$$

Using the $\epsilon_{\text{Tot}}^{\text{OH}}$ and the initial $\delta^{13}\text{C}$ for gas-phase toluene measured, Eq. (2) allows prediction of $\delta^{13}\text{C}_{\text{Allprod}}$ as a function of $c_{\text{Tot}}/\theta c_{\text{Tot}}$ (Fig. 3). For comparison the measured $\delta^{13}\text{C}$ for the different experiment series are included in Fig. 3.

Based on a simple mass balance for all toluene oxidation products and the known POM yields, we

Table 5

Mass balance for secondary POM, sum of all toluene oxidation products, and toluene oxidation products remaining in the gas phase

Experiment	Toluene processing (%)	POM yield (%)	$\delta^{13}\text{C}_{\text{VPDB}}$ POM (‰)	$\delta^{13}\text{C}_{\text{VPDB}}$ all products (‰)	$\delta^{13}\text{C}_{\text{VPDB}}$ gas products (‰)
Series A	15 ± 3	8 ± 2	-32.37 ± 0.08	-32.0 ± 0.3	-32.0 ± 0.3
Series B	7 ± 1	0.4 ± 0.1	-32.90 ± 0.09	-32.3 ± 0.3	-32.3 ± 0.3
Series C	20 ± 2	5 ± 1	-32.57 ± 0.20	-31.9 ± 0.2	-31.8 ± 0.3
Series D	29 ± 2	4 ± 1	-32.19 ± 0.04	-31.6 ± 0.2	-31.5 ± 0.2

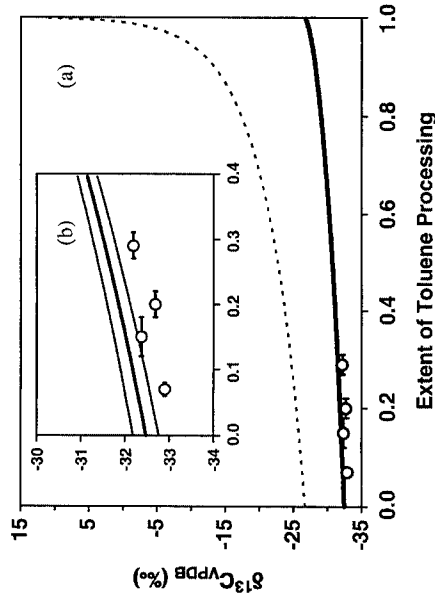


Fig. 3. (a) Dependence of isotope ratio for sum of all products of the oxidation of toluene by OH-radicals on the extent of toluene processing (thick solid line). (b) Expanded scale with standard error (thin solid line). For comparison, the observed $\delta^{13}\text{C}$ values for secondary POM (blank circle) and toluene (dotted line) are also shown.

can calculate the $\delta^{13}\text{C}$ of the toluene reaction products remaining in the gas phase from the measured $\delta^{13}\text{C}$ of POM. The results are summarized in Table 5. The difference between the predicted $\delta^{13}\text{C}$ for all toluene oxidation products and the measured POM $\delta^{13}\text{C}$ ranges from 0.4‰ to 0.7‰, with an average of $0.6\% \pm 0.1\%$ (standard deviation). It should be noted that the largest uncertainty in the predicted $\delta^{13}\text{C}$ for the toluene oxidation products is the uncertainty in the KIE for the reaction of toluene with OH, 0.28‰. The resulting uncertainty of the predicted $\delta^{13}\text{C}$ is slightly below 0.3‰. This systematic uncertainty is not reflected in the variability of the difference in $\delta^{13}\text{C}$ between POM and total oxidation products. Therefore, the overall uncertainty in the difference between total toluene oxidation products and POM is approximately 0.3‰. This indicates that the $\delta^{13}\text{C}$ of POM formed by the gas-phase oxidation of toluene is $0.6\% \pm 0.3\%$ lower than that of the sum of all toluene oxidation products. Due to this small fractionation between POM and the total oxidation products as well as the low yields of POM, it is not surprising that the remaining gas-phase products have effectively the same isotope composition as the sum of all toluene oxidation products.

3. Discussion and conclusions

In our study, POM formed by the OH-radical-induced photochemical oxidation of toluene in the gas phase was found to be between 5.5‰ and 6.2‰

lighter than the precursor compound, toluene. This is only marginally different from the fractionation predicted for the total of all products calculated from the kinetic isotope effect reported by Anderson et al. (2004) for the gas-phase reaction of toluene with the OH-radical. This implies that on average the fractionation between gas-phase and particle-phase products is very small (between 0.4‰ and 0.7‰). It is well established that formation of POM proceeds via reaction chains consisting of several steps; therefore, it is somewhat surprising that the average isotope fractionation of the secondary POM seems to be predominantly determined by the initial reaction step. It is very unlikely that this is just a coincidence, as this finding does not depend on the degree of toluene processing or on the POM yield. However, due to the absence of other studies on $\delta^{13}\text{C}$ of secondary POM we cannot exclude the possibility that the finding is specific for the reaction conditions chosen for our studies.

Based on published observations on substantial gas-phase carbon kinetic isotope effects for reactions of unsaturated organic compounds (Rudolph et al., 2002; Iannone et al., 2003; Anderson et al., 2004), we can rule out the possibility that the reaction steps following the initial reaction of toluene, which lead to POM formation, are not subject to kinetic isotope effects. This leaves in principle two possible explanations for the small fractionation between gas-phase and particle-phase toluene oxidation products. One possibility is that the subsequent reactions of the first-generation products, which lead to POM formation, are complete. Since for complete reaction the isotopic compositions of the reactants are preserved in the products, this will result in POM which has the $\delta^{13}\text{C}$ of the first-generation products. The other possibility is that different isotope fractionation effects associated with the various reaction pathways following the attack of an OH-radical on toluene, leading to POM formation, cancel each other. In the absence of compound-specific isotopic composition information, we cannot distinguish between these possibilities.

Regardless of the possible reasons why the $\delta^{13}\text{C}$ of the secondary POM primarily reflects the isotope effect of the initial reaction step, our finding points towards a promising possibility of using isotope ratio studies to test the applicability of results of laboratory studies for atmospheric conditions. Our results indicate that there is substantial $\delta^{13}\text{C}$ fractionation between the reactant, toluene and

secondary POM. Rudolph and Czuba (2000) have shown that there will be a systematic change in $\delta^{13}\text{C}$ of atmospheric non-methane hydrocarbons depending on the extent of atmospheric processing; therefore, we also expect a systematic dependence between the extent of atmospheric processing of POM precursors and the isotope ratio of secondary POM. However, we are aware that the presently existing information on $\delta^{13}\text{C}$ of secondary POM is insufficient for direct application to the atmosphere. First, as a consequence of the specific requirements for POM formation in a flow reactor, our studies cover only a very limited range of conditions, which significantly differ from atmospheric conditions. Second, total POM in the atmosphere consists of a very complex mixture derived from a variety of sources and atmospheric processes. Therefore compound-specific measurements will be required to study relationships between POM precursors and secondary POM.

Nevertheless, our results show that isotope ratio measurements of secondary POM are a promising tool to test the applicability of laboratory studies for POM to atmospheric conditions as well as to gain deeper insight into the formation mechanisms of secondary POM. Our results demonstrate that under defined conditions, the stable carbon isotope ratios of secondary POM are very reproducible and that the fractionation effects are of a sufficient magnitude to be studied using state-of-the-art compound-specific isotope ratio measurement techniques such as gas chromatography in combination with on-line combustion and isotope ratio mass spectrometry.

Acknowledgements

The authors thank Darrell Ernst and Alina Chivulescu from the Meteorological Service of Canada for the technical support for stable carbon isotope measurements. This project was supported by the Ozone Annex Program in Environment Canada, the Canadian Foundation for Climate and Atmospheric Sciences, and the Natural Sciences and Engineering Research Council of Canada.

References

- Anderson, T.L., Charlson, R.J., Schwartz, S.E., Knutti, R., Boucher, O., Rodhe, H., Heintzenberg, J., 2003. Atmospheric science: climate forcing by aerosols—a hazy picture. *Science* 300 (5622), 1103–1104.
- Anderson, R.S., Iannone, R., Thompson, A.E., Rudolph, J., Huang, L., 2004. Carbon kinetic isotope effects in the gas-phase reactions of aromatic hydrocarbons with the OH radical at (296±4)K. *Geophysical Research Letters* 31 (15), L15108/1–L15108/4.
- Blanchard, P., Brook, J.R., Brazal, P., 2002. Chemical characterization of the organic fraction of atmospheric aerosol at two sites in Ontario, Canada. *Journal of Geophysical Research (Atmospheres)* 107 (D21), ICC10/1–ICC10/8.
- Chesselet, R., Fontugne, M., Buat-Menard, P., Ezat, U., Lambert, C.E., 1981. The origin of particulate organic carbon in the marine atmosphere as indicated by its stable carbon isotopic composition. *Geophysical Research Letters* 8 (4), 345–348.
- Conte, M.H., Weber, J.C., 2002. Plant biomarkers in aerosols: recent isotopic discrimination of terrestrial photosynthesis. *Nature* 417 (6889), 639–641.
- Forstner, H.J.L., Flagan, R.C., Seinfeld, J.H., 1997. Secondary organic aerosol from the photooxidation of aromatic hydrocarbons: molecular composition. *Environmental Science and Technology* 31 (5), 1345–1358.
- Harrington, R.R., Poulson, S.R., Drever, J.I., Colberg, P.J.S., Kelly, E.F., 1999. Carbon isotope systematics of monoaromatic hydrocarbons: vaporization and adsorption experiments. *Organic Geochemistry* 30 (8A), 765–775.
- Huang, L., Norman, A.L., Allison, C.E., Francey, R.J., Ernst, D., Chivulescu, A., Higuchi, K., 2003. Traceability maintenance of high precision stable isotope measurement; $\delta^{13}\text{C}$ & $\delta^{18}\text{C}$ of atmospheric CO_2 at MSC: application to the intercomparison program (Alert, Canada) with CSIRO. In: Toru, S., (Ed.), Report of the 11th WMO/IAEA Meeting of Experts on Carbon Dioxide Concentration and Related Tracer Measurement Techniques, Tokyo, Japan, 2004. World Meteorological Organization Global Atmosphere Watch, pp. 8–14.
- Hurley, M.D., Sokolov, O., Wallington, T.J., Takekawa, H., Karasawa, M., Klotz, B., Barnes, I., Becker, K.H., 2001. Organic aerosol formation during the atmospheric degradation of toluene. *Environmental Science and Technology* 35 (7), 1358–1366.
- Iannone, R., Anderson, R.S., Rudolph, J., Huang, L., Ernst, D., 2003. The carbon kinetic isotope effects of ozone-alkene reactions in the gas-phase and the impact of ozone reactions on the stable carbon isotope ratios of alkenes in the atmosphere. *Geophysical Research Letters* 30 (13), 17/1–17/4.
- Izumi, K., Fukuyama, T., 1990. Photochemical aerosol formation from aromatic hydrocarbons in the presence of nitrogen oxides (NO_x). *Atmospheric Environment Part A: General Topics* 24A (6), 1433–1441.
- Kaplan, I.R., Gordon, R.J., 1994. Non-fossil-fuel fine-particle organic carbon aerosols in southern California determined during the Los Angeles Aerosol Characterization and Source Apportionment Study. *Aerosol Science and Technology* 21, 343–359.
- Kleindienst, T.E., Conner, T.S., McIver, C.D., Edney, E.O., 2004. Determination of secondary organic aerosol products from the photooxidation of toluene and their implications in ambient $\text{PM}_{2.5}$. *Journal of Atmospheric Chemistry* 47 (1), 79–100.
- Narukawa, M., Kawamura, K., Takeuchi, N., Nakajima, T., 1999. Distribution of dicarboxylic acids and carbon isotopic

- compositions in aerosols from 1997 Indonesian forest fires. *Geophysical Research Letters* 26 (20), 3101–3104.
- Norman, A.L., Hopper, J.F., Blanchard, P., Ernst, D., Brice, K., Alexandrou, N., Klouda, G., 1999. The stable carbon isotope composition of atmospheric PAHs. *Atmospheric Environment* 33 (17), 2807–2814.
- Noyes, W.A., 1943. *n*-butyl nitrite. *Organic Syntheses* 2, 108–109.
- Okuda, T., Kumata, H., Naraoka, H., Takada, H., 2002. Origin of atmospheric polycyclic aromatic hydrocarbons (PAHs) in Chinese cities solved by compound-specific stable carbon isotopic analyses. *Organic Geochemistry* 33 (12), 1737–1745.
- Ramanathan, V., Cruzen, P.J., Kiehl, J.T., Rosenfeld, D., 2001. Atmosphere aerosols, climate, and the hydrological cycle. *Science* 294 (5549), 2119–2124.
- Rudolph, J., Czuba, E., 2000. On the use of isotopic composition measurements of volatile organic compounds to determine the "photochemical age" of an air mass. *Geophysical Research Letters* 27 (23), 3865–3868.
- Rudolph, J., Czuba, E., Huang, L., 2000. The stable carbon isotope fractionation for reactions of selected hydrocarbons with OH-radicals and its relevance for atmospheric chemistry. *Journal of Geophysical Research (Atmospheres)* 105 (D24), 29329–29346.
- Rudolph, J., Czuba, E., Norman, A.L., Huang, L., Ernst, D., 2002. Stable carbon isotope composition of nonmethane hydrocarbons in emissions from transportation related sources and atmospheric observations in an urban atmosphere. *Atmospheric Environment* 36 (7), 1173–1181.
- Sakugawa, H., Kaplan, I.R., 1995. Stable carbon isotope measurements of atmospheric organic acids in Los Angeles, California. *Geophysical Research Letters* 22 (12), 1509–1512.
- Sato, K., Klotz, B., Hatakeyama, S., Imamura, T., Washizu, Y., Matsumi, Y., Washida, N., 2004. Secondary organic aerosol formation during the photo-oxidation of toluene: dependence on initial hydrocarbon concentration. *Bulletin of the Chemical Society of Japan* 77 (4), 667–671.
- Simoneit, B.R.T., 1997. Compound-specific carbon isotope analyses of individual long-chain alkanes and alkanolic acids in harrmattan aerosols. *Atmospheric Environment* 31 (15), 2225–2233.
- Stroud, C.A., Makar, P.A., Michelangeli, D.V., Mozurkewich, M., Hastie, D.R., Barbu, A., Humble, J., 2004. Simulating organic aerosol formation during the photooxidation of toluene/NO_x mixtures: comparing the equilibrium and kinetic assumption. *Environmental Science and Technology* 38 (5), 1471–1479.
- Thurston, G.D., Ito, K., Hayes, C.G., Bates, D.V., Lippmann, M., 1994. Respiratory hospital admissions and summertime haze air pollution in Toronto, Ontario: consideration of the role of acid aerosols. *Environmental Research* 65 (2), 271–290.
- Urch, B., Brook, J.R., Wasserstein, D., Brook, R.D., Rajagopalan, S., Corey, P., Silverman, F., 2004. Relative contributions of PM_{2.5} chemical constituents to acute arterial vasoconstriction in humans. *Inhalation Toxicology* 16 (6–7), 345–352.
- Wang, S.C., Paulson, S.E., Grosjean, D., Flagan, R.C., Seinfeld, J.H., 1992. Aerosol formation and growth in atmospheric organic/nitrogen oxide (NO_x) systems—I. Outdoor smog chamber studies of C7- and C8-hydrocarbons. *Atmospheric Environment, Part A: General Topics* 26A (3), 403–420.